Solar and Wind Energy Resource Assessment (SWERA)

High Resolution Solar Radiation Assessment for Sri Lanka

Final country report prepared by DLR

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Content

1. Method description (satellite data, *DNI*-method, *GHI*-method)
2. Model output (*GHI*, *DNI*)
3. Comparison with ground measurements (if available)
4. References

Notice

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1 Method description

Satellite Data

The high resolution solar radiation assessment is based on data of the geostationary satellite Meteosat. Due to the location of the participating SWERA countries, data of Meteosat 7 (M-7) for the years 2000, 2001 and 2002 (for Sri Ghana, Kenya and Ethiopia) and data of Meteosat 5 (M-5) for the years 2000, 2002 and 2003 (for Bangladesh, West-China, Nepal and Sri Lanka) are used. M-5 has its position at 0° latitude and 63° East longitude, M-7 is located at an orbit at 0° latitude and 0° longitude. Figure 1 gives the field of view of both satellites which scans the specific area every 30 minutes with a spatial resolution of 5x5 km².

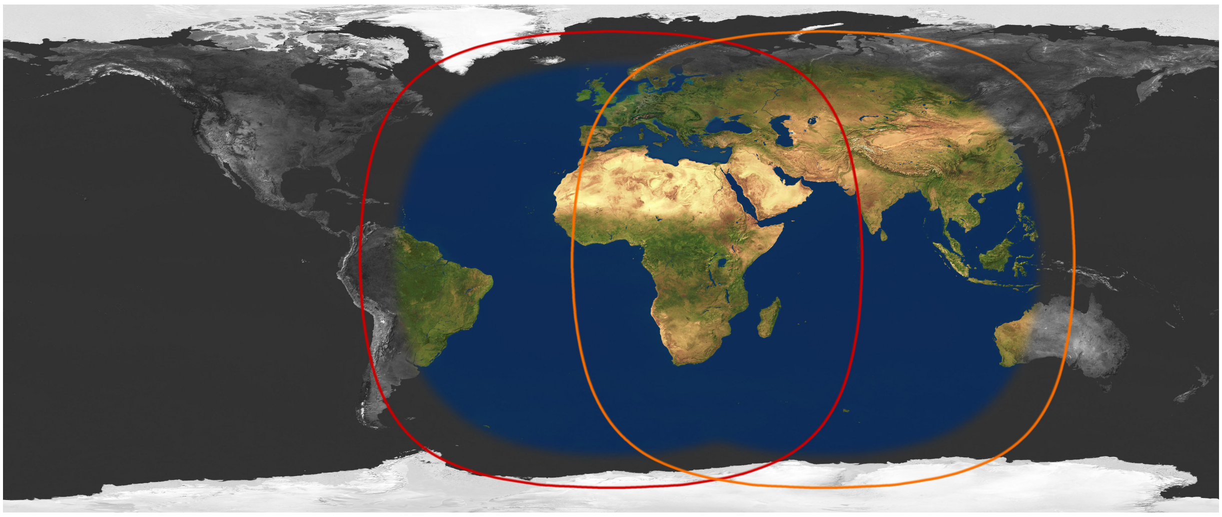


Figure 1: The solar irradiance data is derived from Meteosat a 0° (red circle) and at 63° East (orange circle). The brightened area marks the quantitatively analyzable region. (Meyer et al., 2004).

Data of the visible (VIS) channel, which gives the reflection of the system earth/atmosphere (including clouds) and data of the infrared (IR) channel, which represents the temperature of the surface and atmosphere, are used for gathering information about the clouds. Both are used in a different way to assess the global horizontal (*GHI*) and the direct normal radiation (*DNI*) at ground. Additionally, data of the most important atmospheric components that attenuate the radiation, namely ozone, water vapor and aerosols, are used to take into account the clear-sky conditions of the atmosphere. In the following, the method for deriving *DNI* based on the DLR method and the method for deriving *GHI*, based on a combined method of DLR and SUNY, is described.

Method for Direct Normal Radiation (*DNI*)

The calculation of *DNI* bases on the clear-sky model of Bird and Hulstrom (1981) as described in Iqbal (1983) which was modified by Schillings et al. (2004) for taking into account cloudy conditions with

$$DNI = 0.9751 \cdot I_0 \cdot \tau_R \cdot \tau_{Gas} \cdot \tau_{Ozon} \cdot \tau_{WV} \cdot \tau_{Ae} \cdot \tau_{vis} \cdot \tau_{ir} \quad (1)$$

Each atmospheric transmittance coefficient τ_i is calculated separately using atmospheric input data. All equations for calculating the clear-sky transmittances are described in Iqbal (1983).

Transmittance for Rayleigh scattering

$$\tau_R = \exp\left[-0.0903 m_a^{0.84}\left(1.0 + am_p - am_p^{1.01}\right)\right] \tag{2}$$

Transmittance for equally distributed gas (mainly O_2 and CO_2)

$$\tau_{Gas} = \exp\left(-0.0127 am_p^{0.26}\right) \tag{3}$$

Transmittance for ozone

$$\tau_{Ozon} = 1 - \alpha_{Ozon} \tag{4}$$

$$\alpha_{Ozone} = 0.1611\chi(1.0 + 139.48\chi)^{-0.3035} - 0.002715\chi\left(1.0 + 0.044\chi + 0.0003\chi^2\right)^{-1} \tag{5}$$

$\chi = u \cdot am$, with the vertical ozone layer thickness *u* in cm[NTP] and the airmass *am*.

Transmittance for water vapor

$$\tau_{WV} = 1 - \alpha_{WV} \tag{6}$$

$$\alpha_{WV} = 2.4959\gamma\left[(1.0 + 79.034\gamma)^{0.6828} + 6.385\gamma\right]^{-1} \tag{7}$$

$\gamma = w \cdot am$, with the pressure-corrected relative optical path length of precipitable water *w* in cm[NTP].

Transmittance for aerosols

$$\tau_{Ae} = \exp\left[-k_a^{0.873}\left(1.0 + k_a - k_a^{0.7088}\right)am_p^{0.9108}\right] \tag{8}$$

$$k_a = 0.2758 k_{a\lambda}\big|_{\lambda=0.38\mu m} + 0.35 k_{a\lambda}\big|_{\lambda=0.5\mu m} \tag{9}$$

with the aerosol optical thickness $k_{a\lambda}$ at the wavelength 0.38 μm und 0.5 μm.

Transmittance for clouds

using the visible Cloud-Index *CI_vis*

$$\tau_{vis} = e^{\left(-CI_{vis} \cdot 0.1\right)} \tag{10}$$

and using the infrared Cloud-Index *CI_ir*

$$\tau_{ir} = e^{\left(-CI_{ir} \cdot 0.07\right)} \tag{11}$$

For the clear-sky atmospheric transmittance, the airmass is needed which is calculated by

$$am = \frac{1}{\left[\cos\Theta_Z + 0.15(93.885 - \Theta_Z)\right]^{-1.253}} \tag{12}$$

The pressure correction is made by

$$am_p = am \cdot \frac{p}{1013.25} \tag{13}$$

with

$$\frac{p}{p_0} = \exp(-0.0001184 z) \tag{14}$$

The clear-sky radiation is calculated each 20 minutes (10,30,50 minutes of each hour) for the maps and each 5 minutes (5,10,15,…,55,60 minutes each hour) for the time series. The influence of the clouds is taken into account hourly, therefore all maps (monthly and annual average daily sums) and time series are based on an hourly calculation of the radiation. The DLR -model output for *DNI* is sampled at a 10km spatial resolution.

Method for Global Horizontal Radiation (*GHI*)

The calculation of *GHI* bases on the method of Perez et al (2002) and Ineichen and Perez (2002). *GHI* is calculated with (Perez et al., 2002)

$$GHI = ktm \cdot Ghc \cdot (0.0001 \cdot ktm \cdot Ghc + 0.9) \tag{15}$$

with *ktm*

$$ktm = 2.36 \cdot CI^5 - 6.2 \cdot CI^4 + 6.22 \cdot CI^3 - 2.63 \cdot CI^2 - 0.58 \cdot CI + 1 \tag{16}$$

GHI is calculated using the cloud information based on infrared (IR) and visible (VIS) Meteosat data which lead to a single Cloud-Index *CI* with

$$CI = \max(CI\_vis, CI\_ir) \tag{17}$$

For the determination of the clear-sky global irradiance G_{hc} the new formulation as described in Perez et al (2002) is used with

$$Ghc = cg1 \cdot I_0 \cdot \cos\Theta_Z \cdot \exp(-cg2 \cdot am \cdot (fh1 + fh2 \cdot (TL - 1)))\exp(0.01 \cdot am^{1.8}) \tag{18}$$

with

- *cg1* = (0.0000509 * alt + 0.868)
- *cg2* = (0.0000392 * alt + 0.0387)
- I_0 = Solar constant (eccentricity corrected)
- Θ_z = solar zenith angle
- *fh1* = exp(-alt / 8000)
- *fh2* = exp(-alt / 1250)
- *am* = elevation corrected air mass
- *alt* = altitude in meters
- T_L = Linke turbidity

Due to missing values of the Linke turbidity T_L for the parameterization of the clear-sky atmosphere, data of the atmospheric components ozone, water vapor and aerosols are used. These atmospheric data are also used for the *DNI*. To derive T_L from atmospheric data we use the following formulation as described by Ineichen and Perez (2002) with

$$Tl = ((11.1 \cdot \ln(b \cdot \frac{I_0}{B_{ncl}})) / am) + 1 \tag{19}$$

with $b = 0.664 + (0.163 / fh1)$ (20)

and the clear-sky direct normal irradiance B_{ncl}

$$B_{ncl} = I_0 \cdot \tau_{\mathrm{ra}} \cdot \tau_{\mathrm{ae}} \cdot \tau_{\mathrm{o3}} \cdot \tau_{\mathrm{ga}} \cdot \tau_{\mathrm{wv}} \tag{21}$$

The calculation of transmittance coefficients τ_i and the used atmospheric input data are described in the method for the *DNI*.

The clear-sky radiation is calculated each 20 minutes (10,30,50 minutes of each hour) for the maps and each 5 minutes (5,10,15,…,55,60 minutes each hour) for the time series. The influence of the clouds is taken into account hourly, therefore all maps (monthly and annual average daily sums) and time series are based on an hourly calculation of the radiation. The DLR/SUNY-model output for *GHI* is sampled at a 10km spatial resolution.

Input Data

Elevation

For the airmass pressure correction, the elevation from the GLOBE database from the USGS U.S. Geological Survey [http://rockyweb.cr.usgs.gov/elevation/dpi_dem.html] is used, (Hastings and Dunbar, 1998).

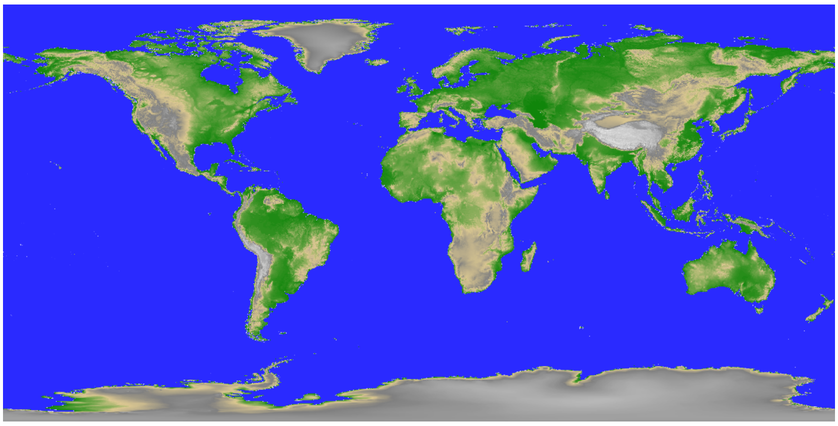


Figure 2: Elevation from GLOBE.

Ozone

The monthly ozone data are taken from TOMS published by the NASA/GSFC TOMS Ozone Processing Team [http://toms.gsfc.nasa.gov/], (McPeters et al., 1998).

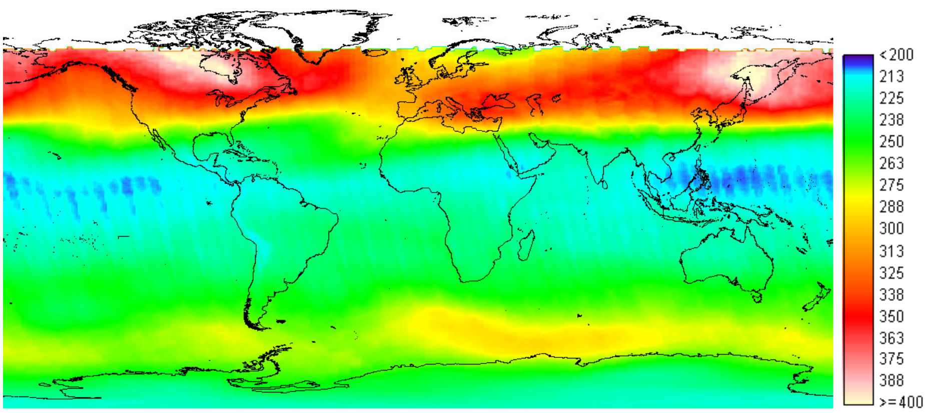


Figure 3: Ozon monthly average for February 2003 in [DU] from TOMS

Water vapor

The daily water vapor data are taken from the NOAA-CIRES Climate Diagnostics Center in Boulder Colorado, USA (NCEP/NCAR) [http://www.cdc.noaa.gov/] (Kalnay et al., 1996).

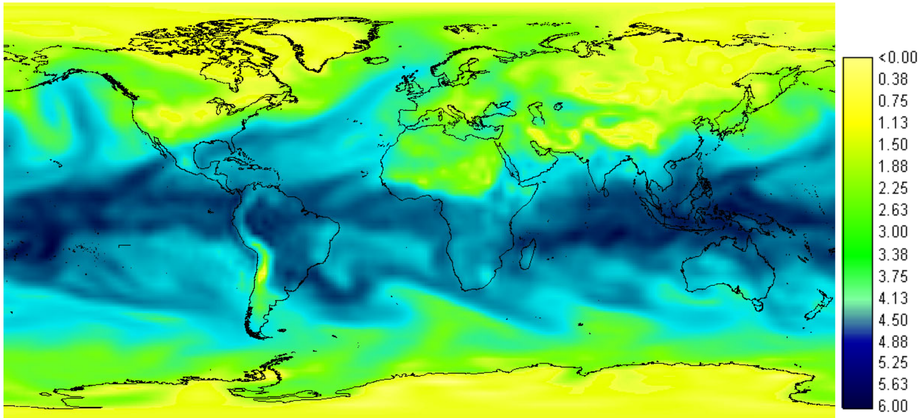


Figure 4: Water vapor daily mean for 7. February 2003 in cm[NTP] from NCEP/ NCAR-Reanalysis

<u>Aerosol</u>

The monthly climatological aerosol optical thickness data are taken from NASA-GACP, [http://gacp.giss.nasa.gov/index.html], (Mishchenko et al, 2002).

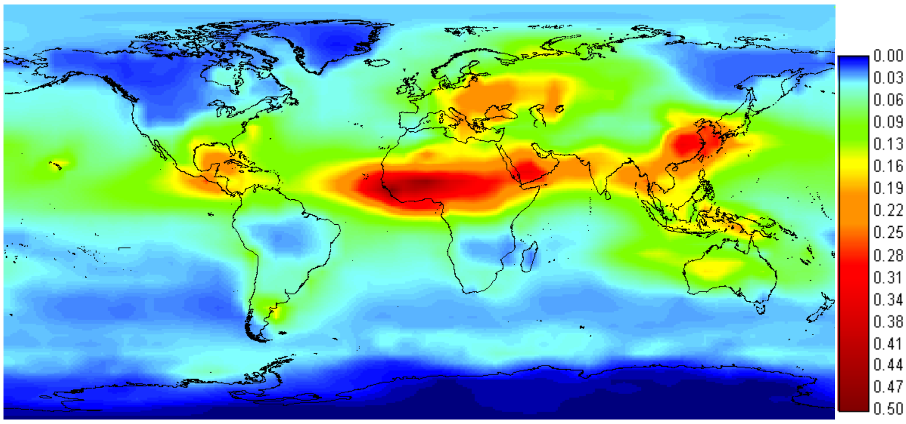


Figure 5: Aerosol optical thickness for February from NASA-GACP.

<u>Clouds</u>

The hourly cloud information are based on half-hourly Meteosat-5 IR and VIS data (© EUMETSAT, 2004). The determination of the cloud indices is described in detail in Mannstein et al. (1999) and Schillings et al. (2004). The basic approach for deriving VIS cloud information is described with

$$CI\_vis = \frac{\rho - \rho_{\min}}{\rho_{\max} - \rho_{\min}} \tag{22}$$

where ρ is the actual reflectivity measured by the satellite, ρ_{min} corresponds to the surface albedo and ρ_{max} is the maximum reflectivity measured for overcast cloudy conditions. The similar approach is used for IR-data, with the actual, minimum and maximum brightness temperatures T measured by the satellite:

$$CI\_ir = \frac{T_{\min} - T}{T_{\min} - T_{\max}} \tag{23}$$

Figure 6: Field of view of Meteosat (M-7, IR-channel) © 2004, EUMETSAT.

2 Model output

The solar radiation is calculated for the complete country for the years 2000, 2001 and 2002. The data are made available in a digital GIS-format (ESRI Vector-Shapefile). Within this report, maps of the annual average daily total sum of *GHI* and *DNI* are presented. The complete database (ESRI-Shapefile and MS-Access database) can be downloaded from the SWERA-homepage (http://swera.unep.net). Within the ESRI Vector-Shapefile, 3 annual and 36 monthly average daily total sums of *GHI* and *DNI* are given for each 10km x 10km georeferenced pixel as shown in the following figures. Additional, hourly time series for the same time period for several interesting sites are delivered in a separate ASCII-File. The output time of the hourly data is UTC.

Time series

For following sites hourly time series of GHI and DNI for three years are calculated:

Stations/Sites	Lat(degree)	Long(degree)	Elevation (m)
Colombo	6.91	79.85	10
Jaffna	9.68	79.98	10
Kalmunai	7.40	81.80	10

The hourly time series can be downloaded from the SWERA web-site. All ASCII-files are compressed to one single ZIP-file. The name convention of the ASCII-file name is:

Country_Sitename_Lat_Lon_Z_Year.dat

for example: Sri_Lanka_Colombo_N6.91_E79.85_Z10_2000.dat

Important notice: The following maps show classified values of *kWh/m²/day* with a common color ramp for all SWERA countries to give a first impression of the solar regime for each country and for easier comparison with other countries. The provided digital GIS data (available at http://swera.unep.net) give the real (and not classified!) values in *Wh/m²/day* for each georeferenced pixel with a signal resolution of 1 Wh/m²/day.

Global Horizontal Radiation

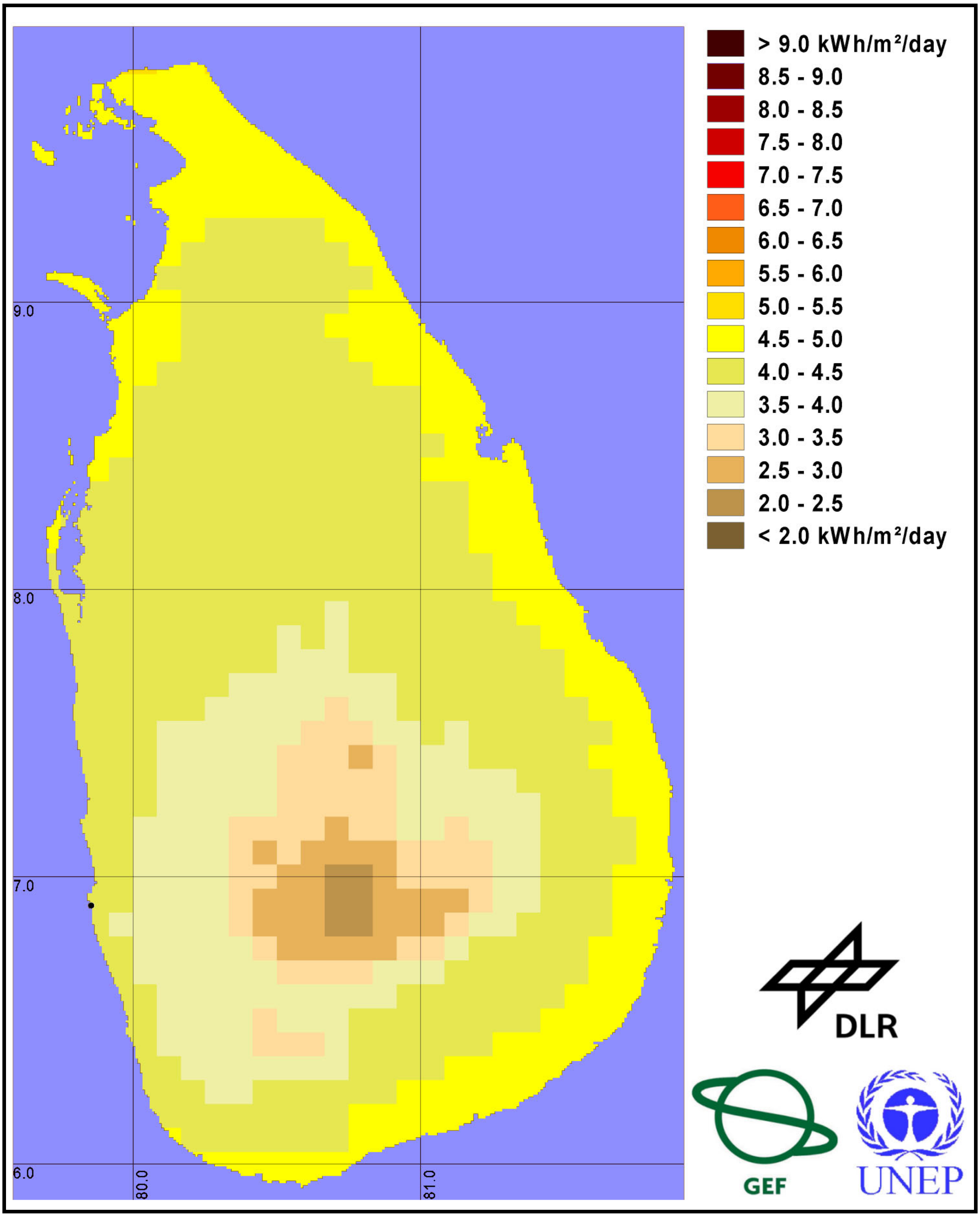


Figure 7: Annual average daily total sum of *GHI* kWh/m²/day for Sri Lanka 2000

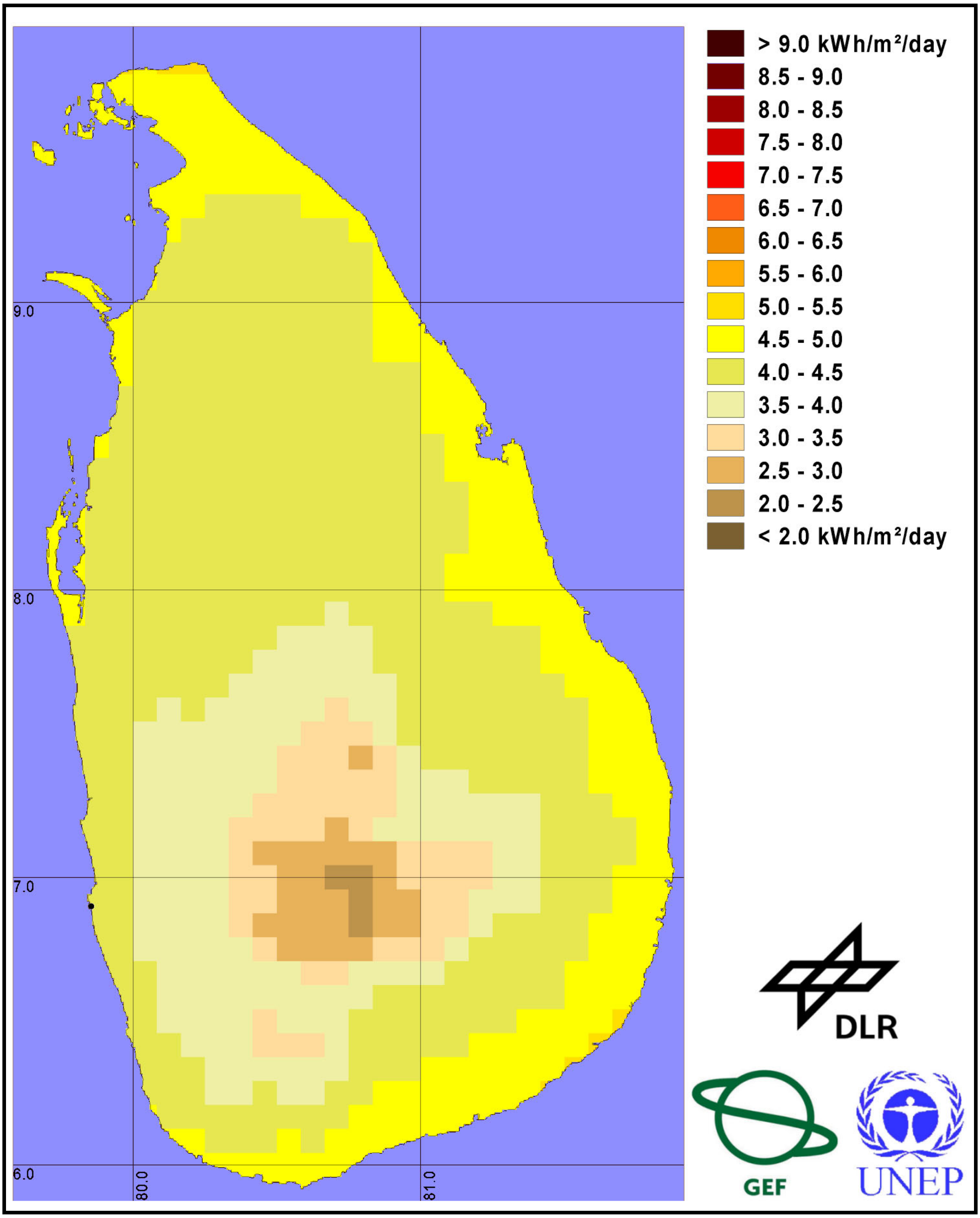


Figure 8: Annual average daily total sum of *GHI* in kWh/m²/day for Sri Lanka 2002

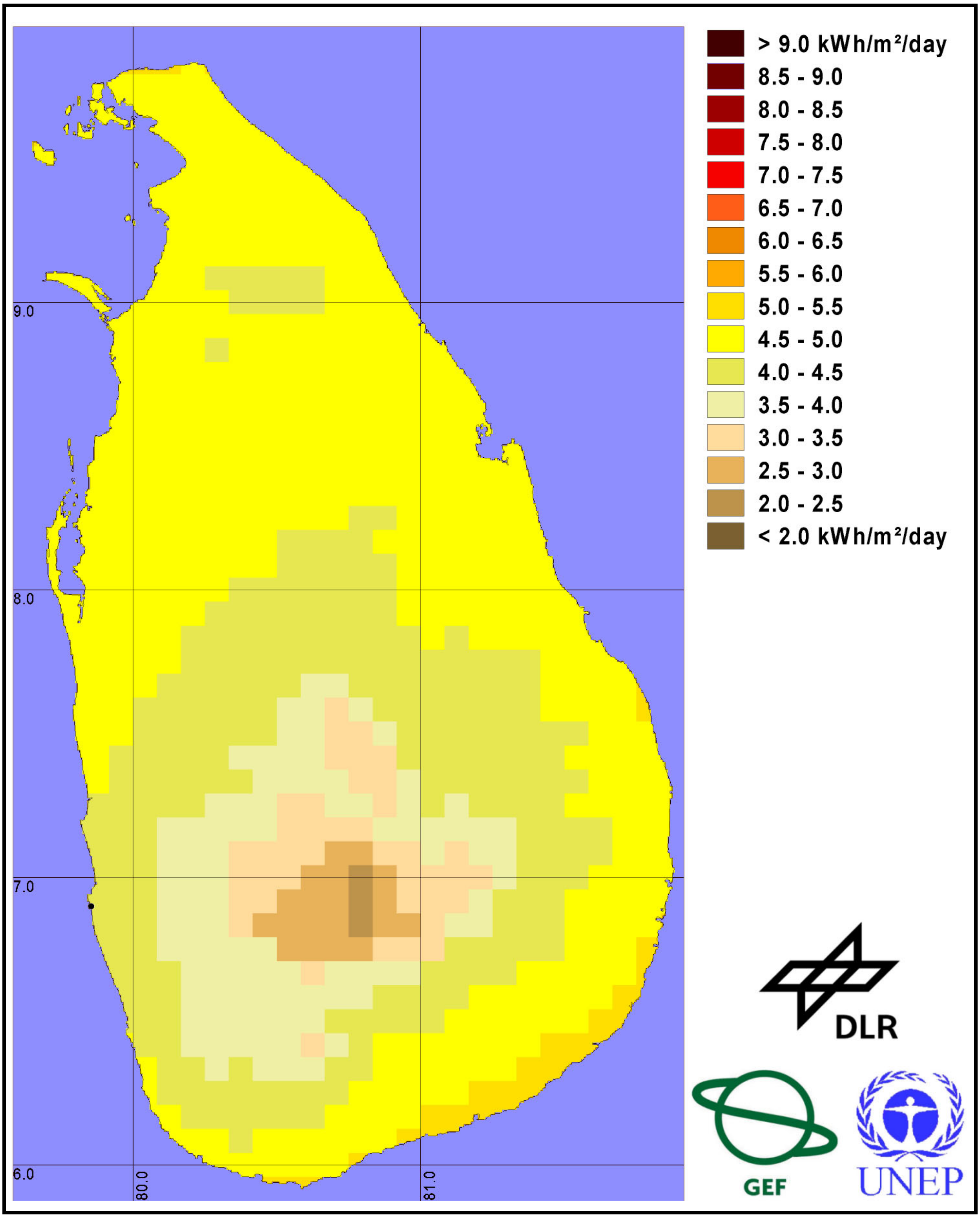


Figure 9: Annual average daily total sum of *GHI* in kWh/m²/day for Sri Lanka 2003

Direct Normal Radiation

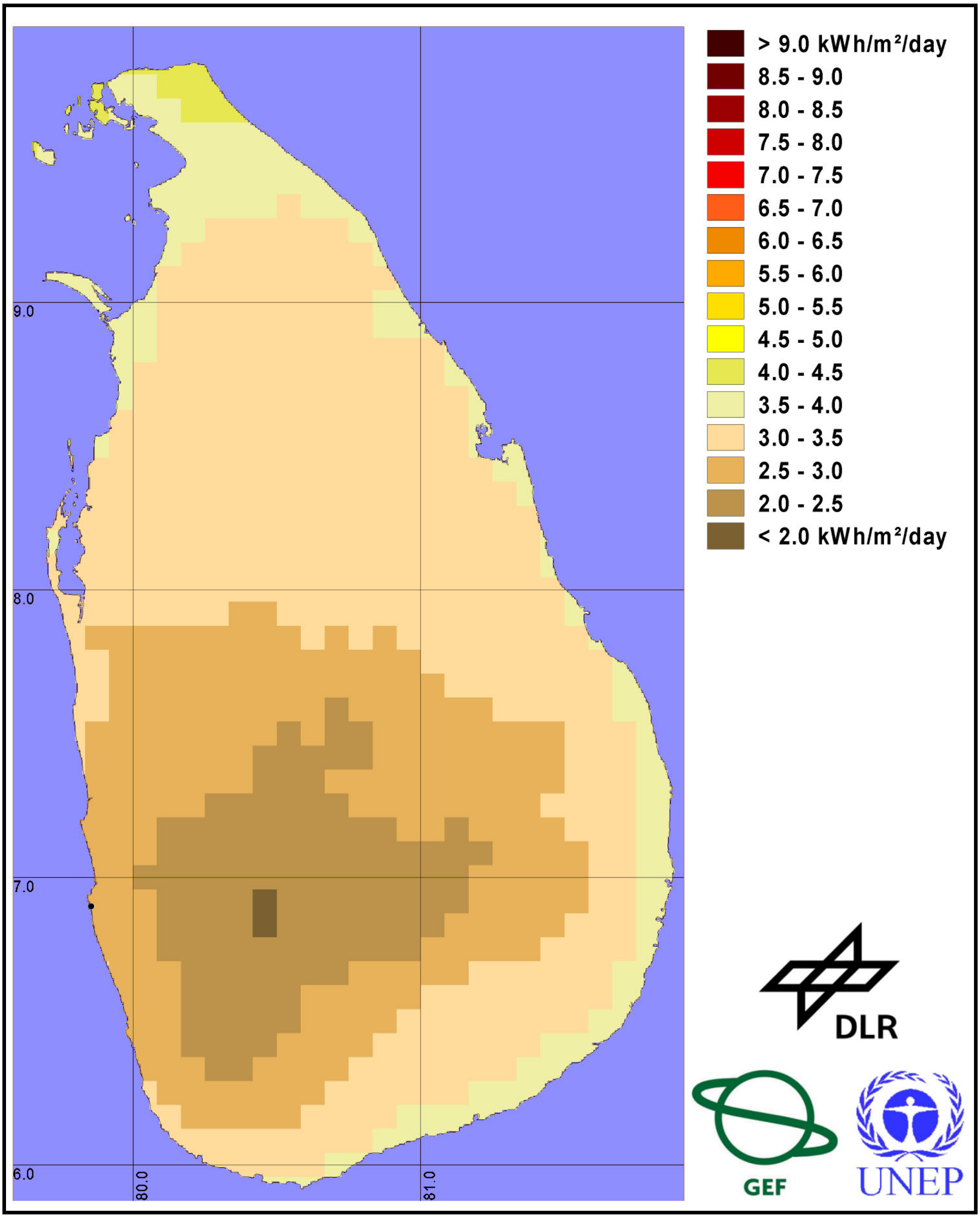


Figure 10: Annual average daily total sum of *DNI* in kWh/m²/day for Sri Lanka 2000

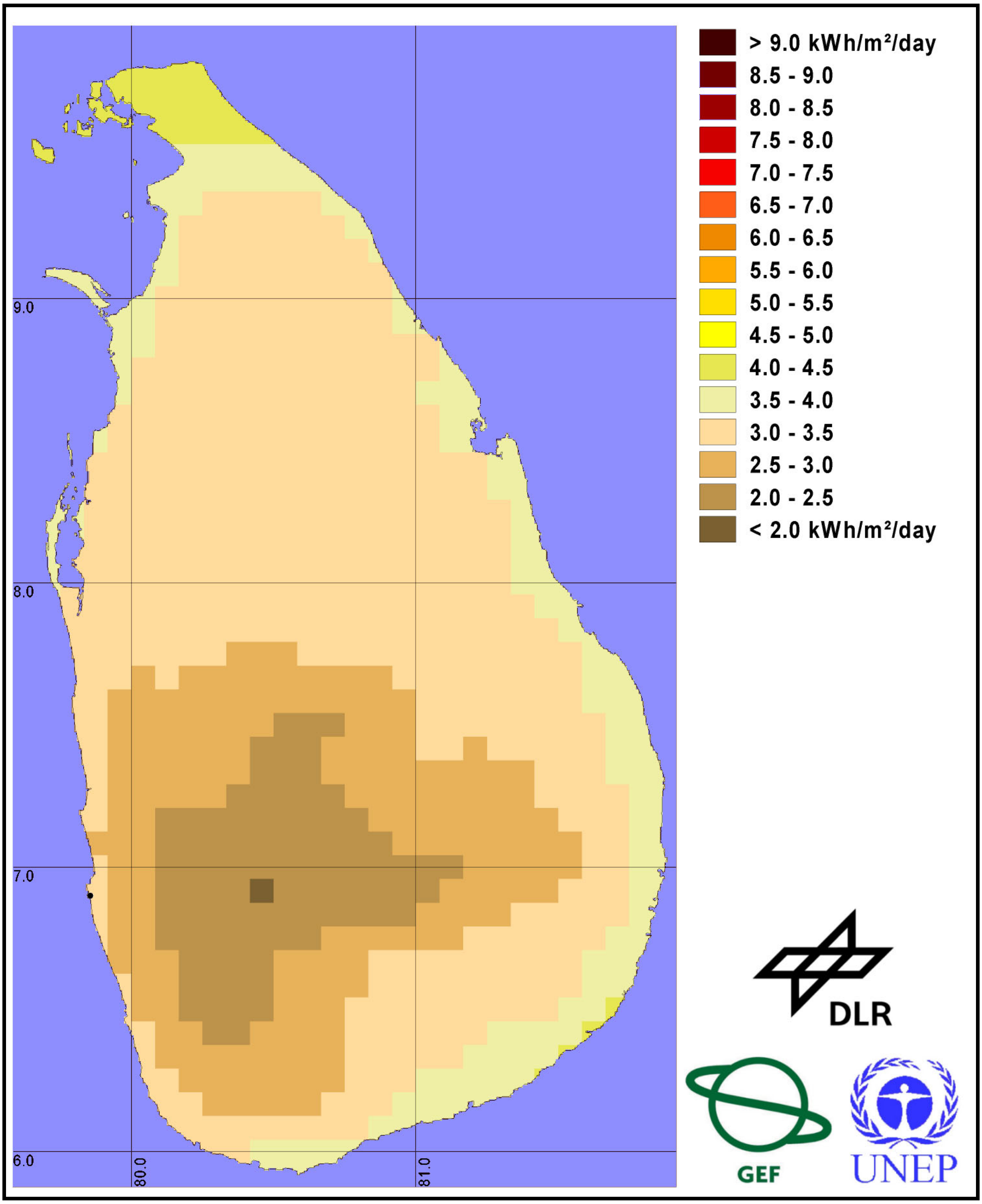


Figure 11: Annual average daily total sum of *DNI* in kWh/m²/day for Sri Lanka 2002

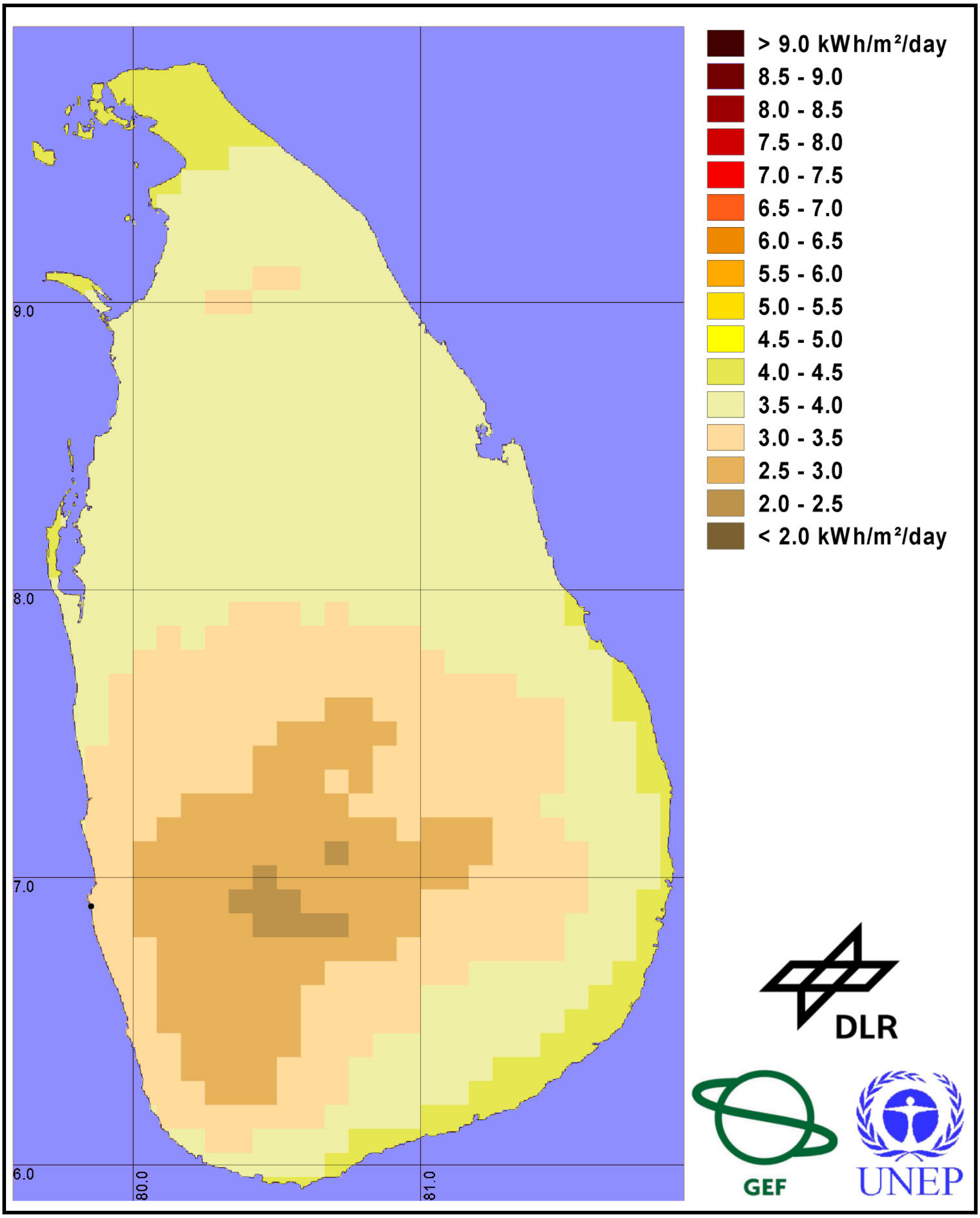


Figure 12: Annual average daily total sum of *DNI* in kWh/m²/day for Sri Lanka 2003

3 Comparison with ground measurement in Sri Lanka

Ground data: Ground measurements of the global horizontal irradiance are available for one site in Sri Lanka (Colombo, Latitude: 6 54' 30", Longitude: 79 51' 30", Elevation: 10m) as shown in figure 13. The data are provided by the Renewable Energy Department of the NERD Centre of Sri Lanka. The *GHI* is measured by a high precision Pyranometer OSK 15306. The measured values are integrated to hourly values for this comparison. Hourly values are available for the period 1998 to 2003. For the comparison values of the year 2000, 2002 and 2003 are taken. All ground data are provided from NERDC within the SWERA-project.

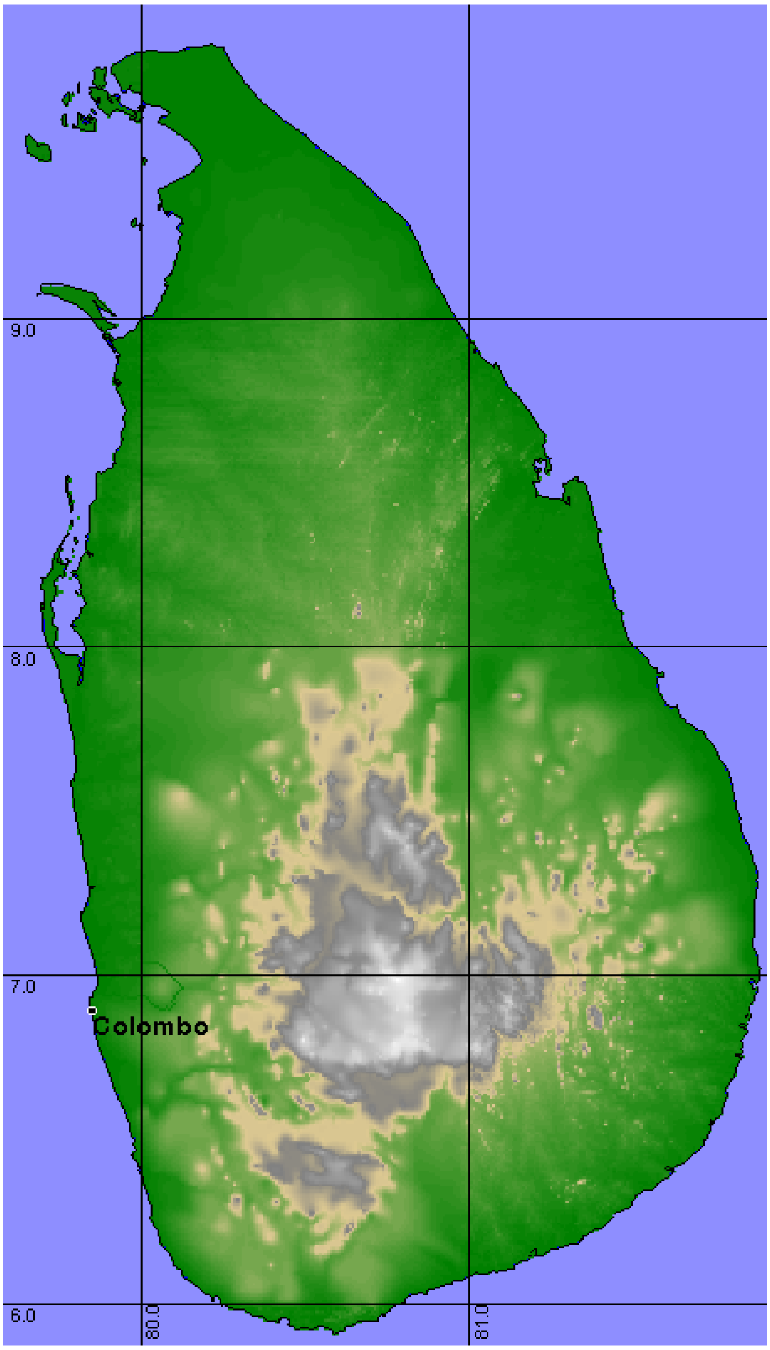


Figure 13: Measurement site in Sri Lanka.

Satellite data: The complete years 2000, 2002 and 2003 are calculated and there are no data missing.

Comparison results: Several comparisons are performed for the site Colombo. The averaged hourly values of the years 2000, 2002 and 2003 are compared. The relative Mean Bias Deviation (rMBD) for the hourly comparison is 1.03%, the relative Root Mean Square Deviation (rRMSD) 37.11%. 4662 averaged hourly values are compared. The result is also shown in figure 14.

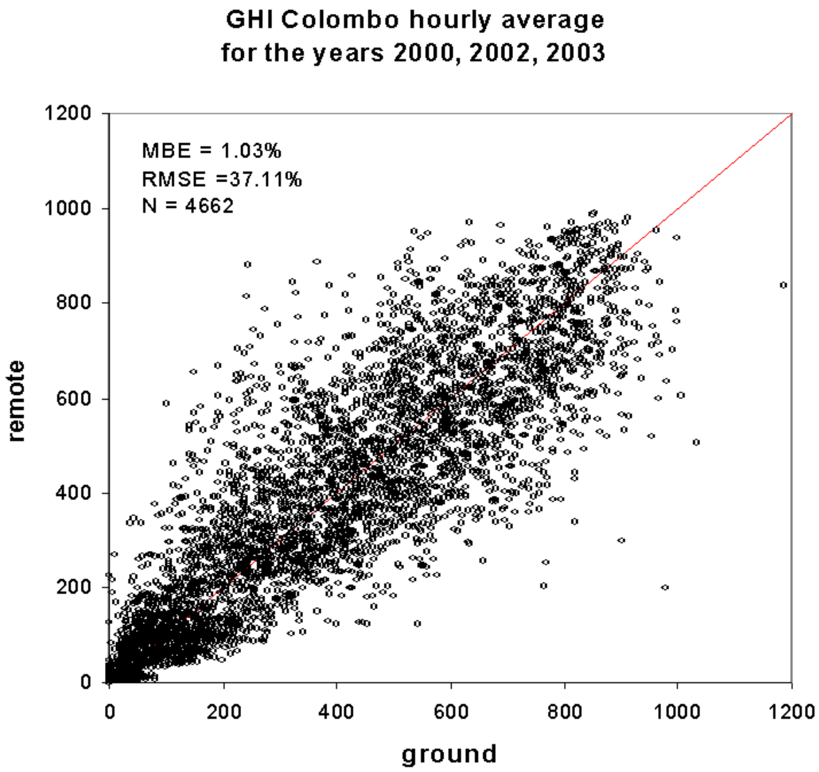


Figure 14: Comparison of ground-measured and satellite-derived hourly values. The hourly values are averaged for the three years 2000, 2002 and 2003.

The next figure shows the monthly average daily sum of *GHI* (ground and satellite), for each year and for the 3-years average (2000, 2002, 2003) for the site Colombo. The rMBD for the monthly values (not averages!) is 2.45% and the rRMSD 9.35%.

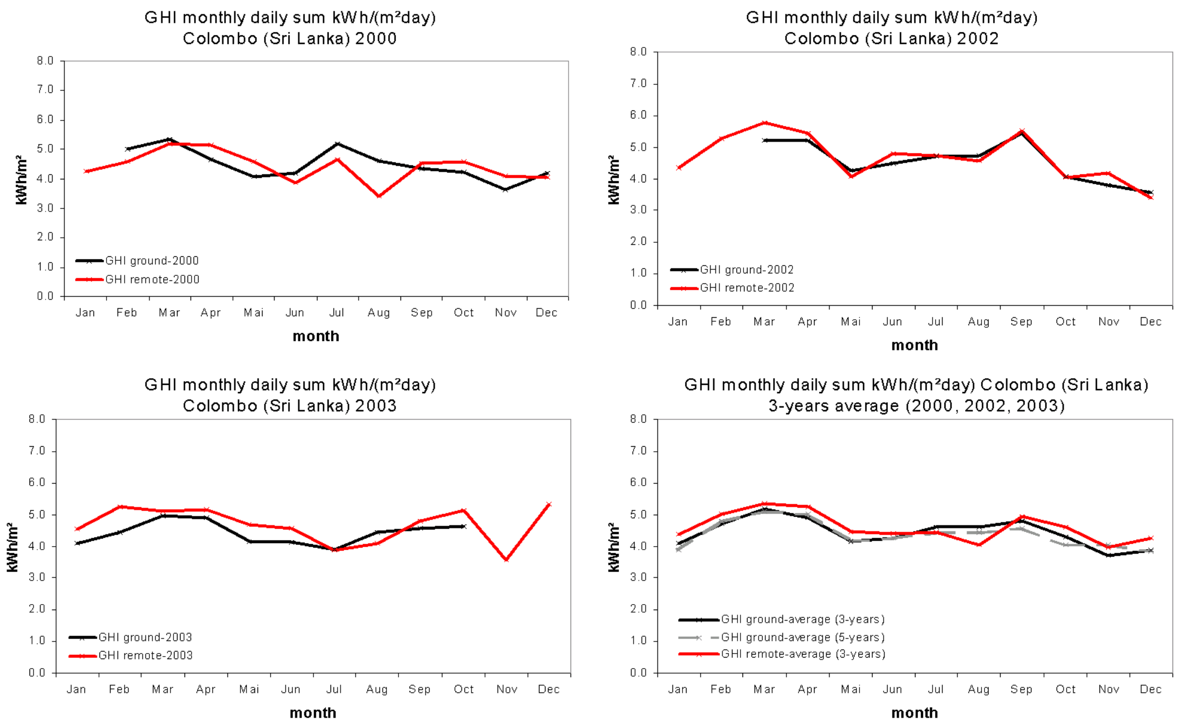
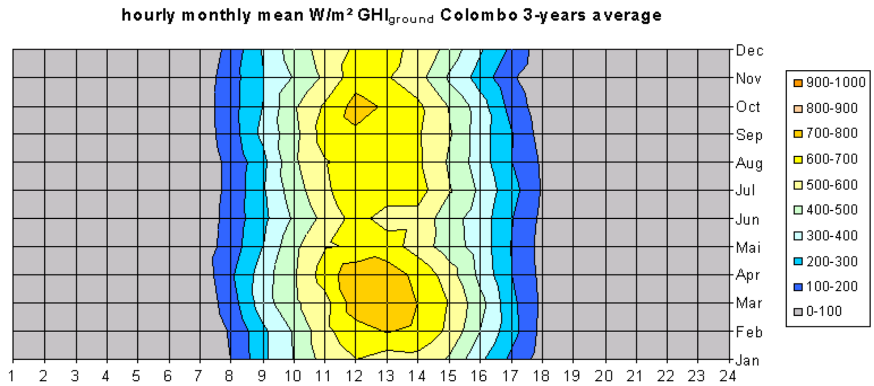


Figure 15: Monthly average daily sum of *GHI* (ground and satellite), for each year and for the 3-years average (2000, 2002, 2003) for the site Colombo in kWh/m²day. The red line represents the satellite-derived data, the black one the ground-measured data. For the averaged values (bottom, right) the 5-years average of the ground data for the period 1999-2003 is also shown (dashed grey line).

To give an impression of the global radiation region during on year, the hourly monthly mean GHI in W/m² is shown in the following figures.



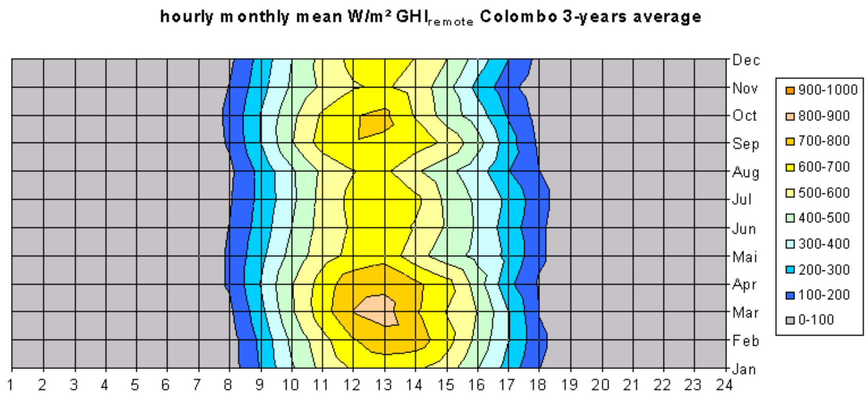


Figure 15: hourly monthly mean values in W/m² for the ground-measured (top) and the satellite-derived (down) data. The 3-years average values are shown.

Acknowledgement:

All Meteosat data is under copyright of EUMETSAT (© 2004), Darmstadt, Germany. Many thanks for access to the data set go to the crew of the MARF (Meteosat Archive and Retrieval Facility, Darmstadt) and to our colleagues from DLR-DFD (Deutsches Fernerkundungs-Datenzentrum). Data on aerosol is provided by NASA-GACP Global Aerosol Climatology Project. We also acknowledge the use of the water vapor data from the NCEP Reanalysis by NOAA-CIRES Climate Diagnostics Center, Boulder, Colorado, US and the use of the TOMS ozone data provided by the NASA-Goddard Space Flight Center (GSFC), Washington, DC, USA. Special thanks to Rüdiger Buell and Hermann Mannstein at the Institut für Physik der Atmosphäre for doing the archive logistics and cloud data processing of the Meteosat 5 and Meteosat 7 data. Thank you to Mahinsasa Narayana from NERDC for preparing and providing the ground measurements of Colombo.

4 References

Bird R.E. and R.L. Hulstrom (1981): Review, evaluation and improvement of direct irradiance models. *J. of Sol. Energy Engineering* **103**, 182-192.

Hastings D.A. and P.K. Dunbar (1998): Development & Assessment of the Global Land One-km Base Elevation Digital Elevation Model (GLOBE). *International Society of Photogrammetry and Remote Sensing (ISPRS) Archives* **32** (4), 218-221. (http://www.ngdc.noaa.gov/seg/topo/globe.shtml)

Ineichen P. and R. Perez (2002): A new airmass independent formulation for the Linke turbidity coefficient. *Solar Energy*, **73** (3), 151-157.

Iqbal M. (1983): An introduction to solar radiation. Academic Press, Toronto.

Kalnay E., M. Kanamitsu, R. Kistler, W. Collins, D. Deaven, L. Gandin, M. Iredell, S. Saha, G. White, J. Woollen, Y. Zhu, A. Leetmaa, R. Reynolds, M. Chelliah, W. Ebisuzaki, W. Higgins, J. Janowiak, K.C. Mo, C. Ropelewski, J. Wang, R. Jenne and D. Joseph (1996): The NMC/NCAR 40-Year Reanalysis Project. Bulletin of the American Meteorological Society **77** (3), 437–472.

Mannstein H., H. Broesamle, C. Schillings and F. Trieb (1999): Using a METEOSAT cloud index to model the performance of solar-thermal power stations. *Procs. of Eumetsat Conference Copenhagen*, 239-246.

McPeters R.D. , P. K. Bhartia, A. J. Krueger, J. R. Herman, C. G. Wellemeyer, C. J. Seftor, G. Jaross, O. Torres, L. Moy, G. Labow, W. Byerly, S. L. Taylor, T. Swissler and R. P. Cebula (1998): Earth Probe Total Ozone Mapping Spectrometer (TOMS) Data Products User's Guide. *NASA Technical Publication* 1998-206895.

Meyer R., C. Hoyer, E. Diedrich, C. Schillings, S. Lohmann, M. Schroedter-Homscheidt, R. Buell and F.Trieb (2004): Solar Energy Mining: A High-Quality Satellite-based Service to Provide Direct Solar Radiation for Europe, Brazil, Africa and Asia. Presentation at SolarPaces 2004, Oaxaca, Mexico.

Mishchenko M., J.E. Penner and D. Anderson (Ed.) (2002): Global Aerosol Climatology Project. Special Issue of Global Aerosol Climate, *J. of Atmospheric Science* **59** (3).

Perez R., P. Ineichen, K. Moore, M. Kmiecik, C. Chain, R. George and F. Vignola (2002): A new operational satellite-to-irradiance model - description and validation. *Solar Energy*, **73** (5), 307-317.

Schillings C., H. Mannstein and R. Meyer (2004): Operational method for deriving high resolution direct normal irradiance from satellite data. *Solar Energy* **76**, 475-484.